

Synoptic-Scale Analysis of Freezing Rain Events in Montreal, Quebec

Gina Ressler, Shawn Milrad, Eyad Atallah, and John Gyakum

Department of Atmospheric and Oceanic Sciences, McGill University,
Montreal, Quebec, Canada

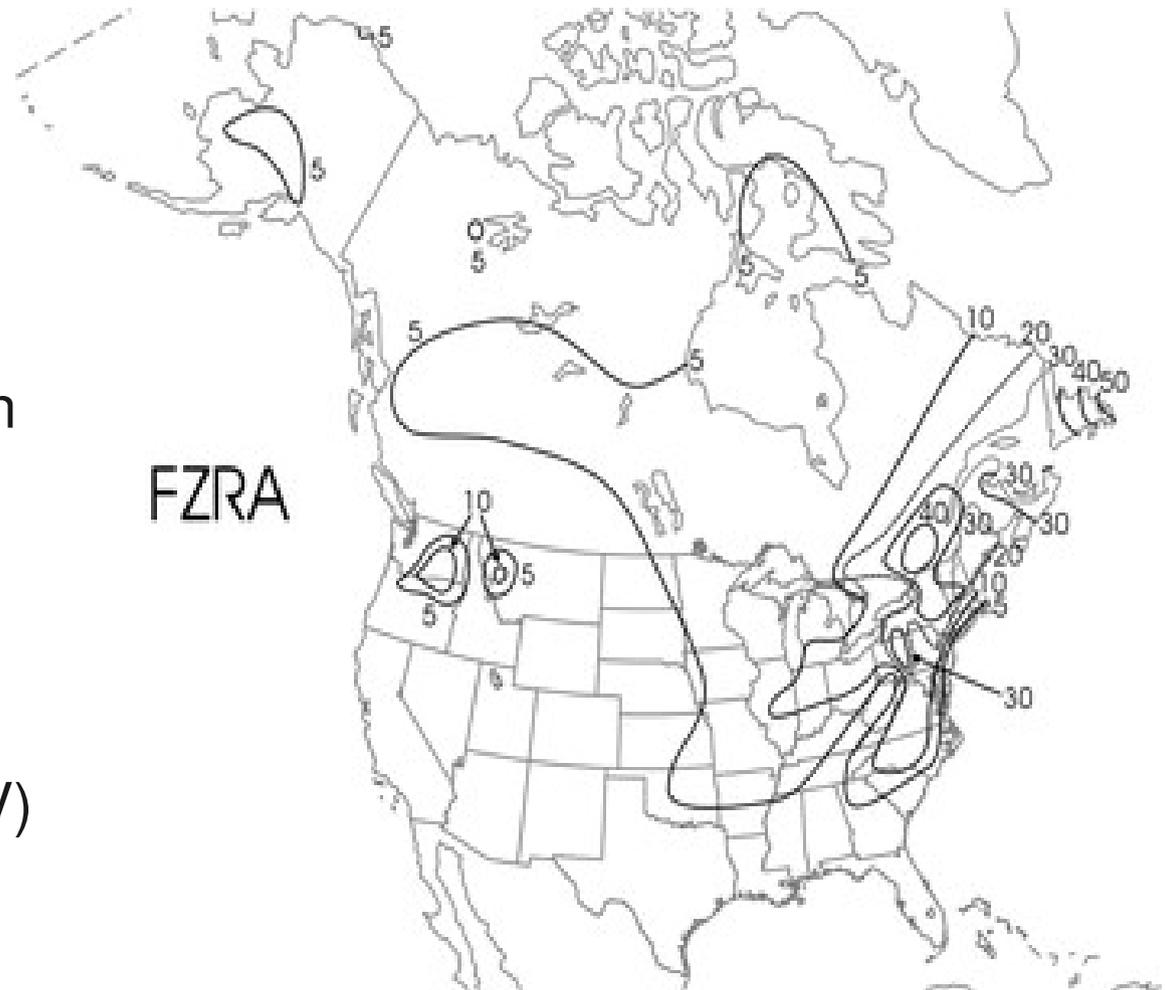


*Great Lakes Operational Meteorology Workshop
Ithaca, New York, USA
March 21-23, 2011*



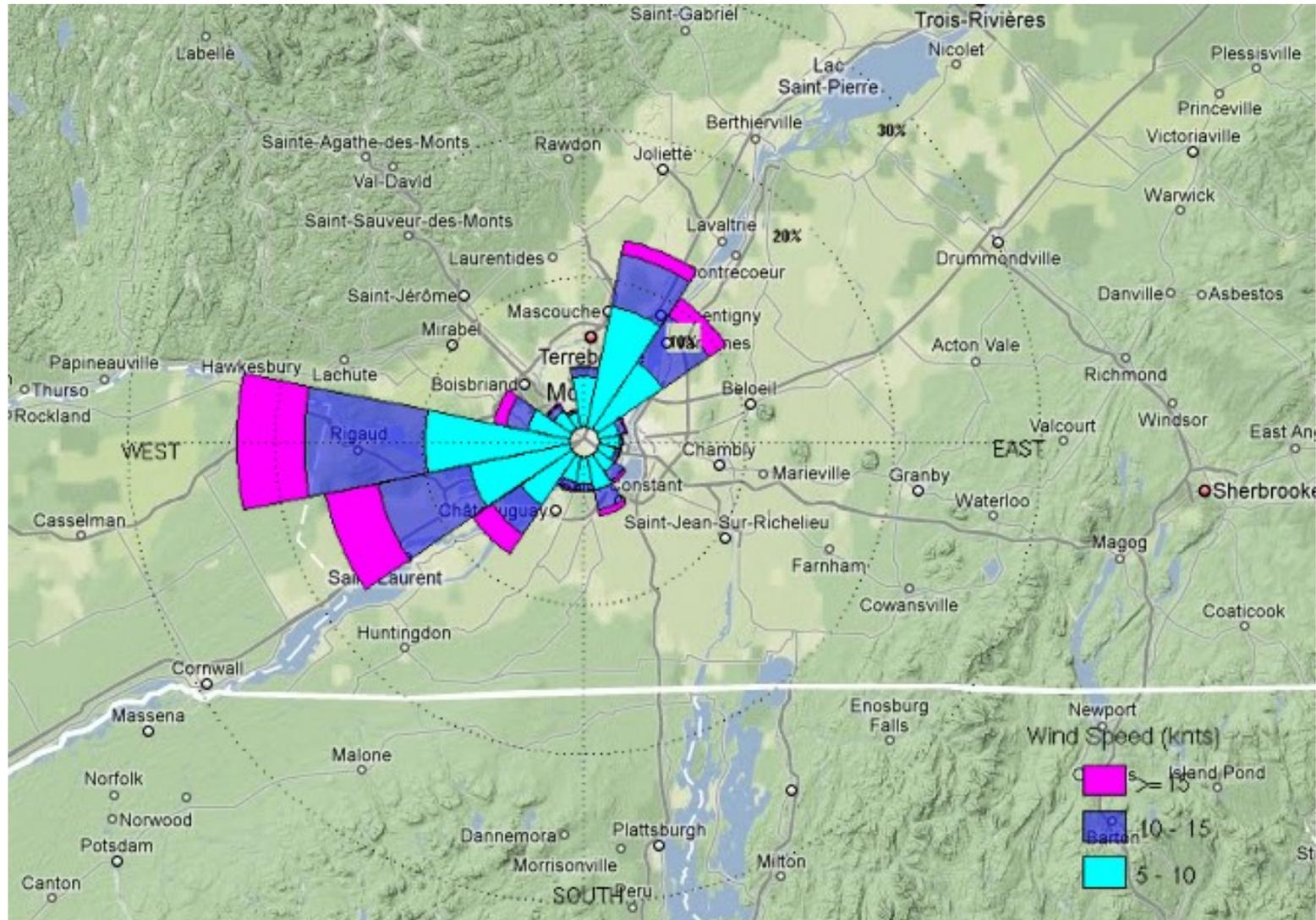
Motivation

- Freezing rain is a major environmental hazard in eastern Canada/US
 - 1998 Ice Storm - \$3 billion in damage, left over one million people without power
- Especially common in the St. Lawrence River Valley (SLRV)
- More insight – better forecasts



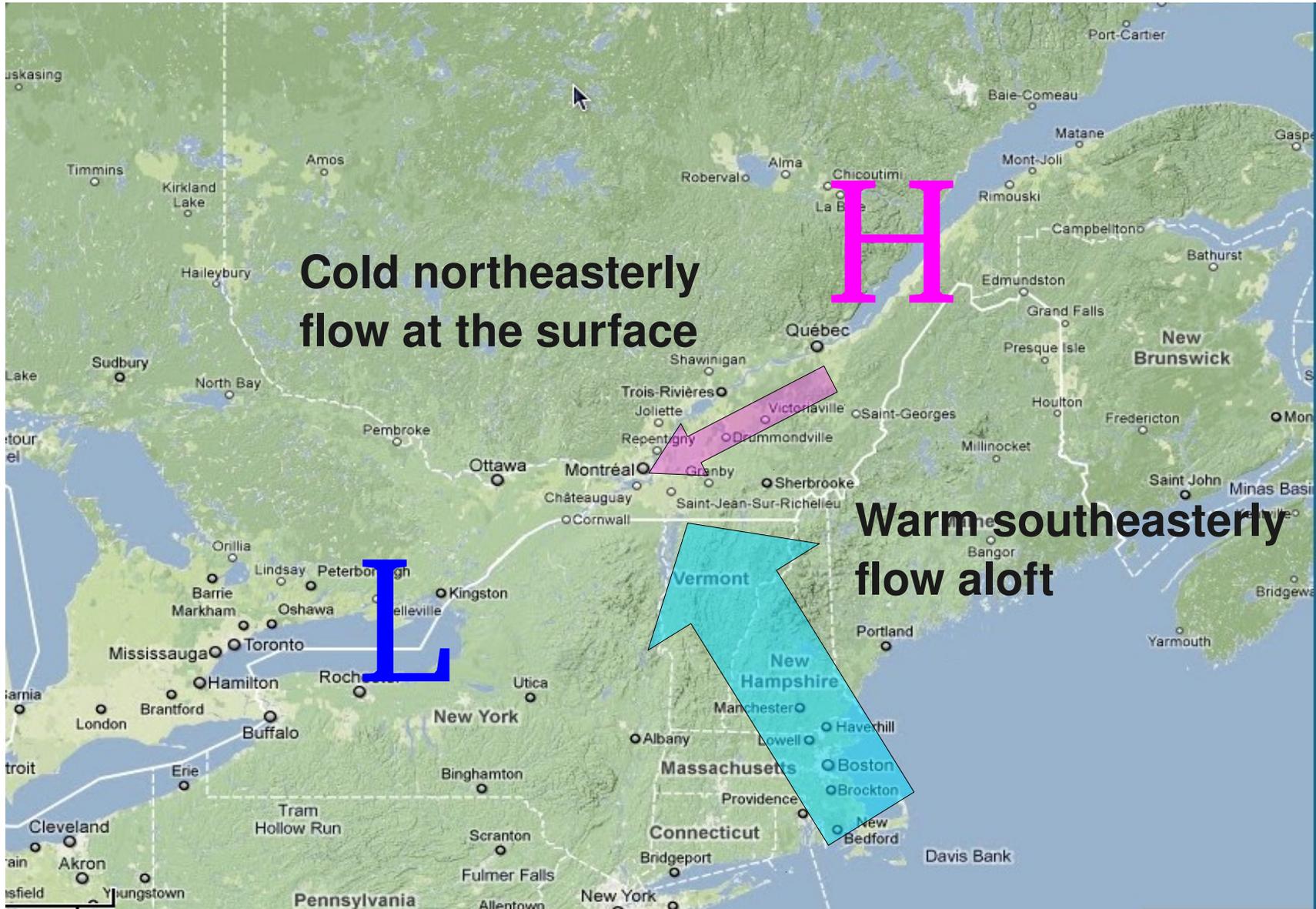
Median annual hours of freezing rain from 1979 to 1990 (Fig. 2, Cortinas et al. 2004)

DJF observed surface winds at Montreal, Quebec (YUL) 1979-2002



Courtesy of Alissa Razy

Weather Scenario for Freezing Rain



Adapted from Milton and Bourque 1999

Objectives

- Our knowledge of freezing rain at Montreal has been derived primarily from case study analyses
- Goals:
 - Construct a complete list of freezing rain events for the time period 1979-2008
 - Identify key synoptic-scale features of typical Montreal freezing rain events



Montreal, December 2008

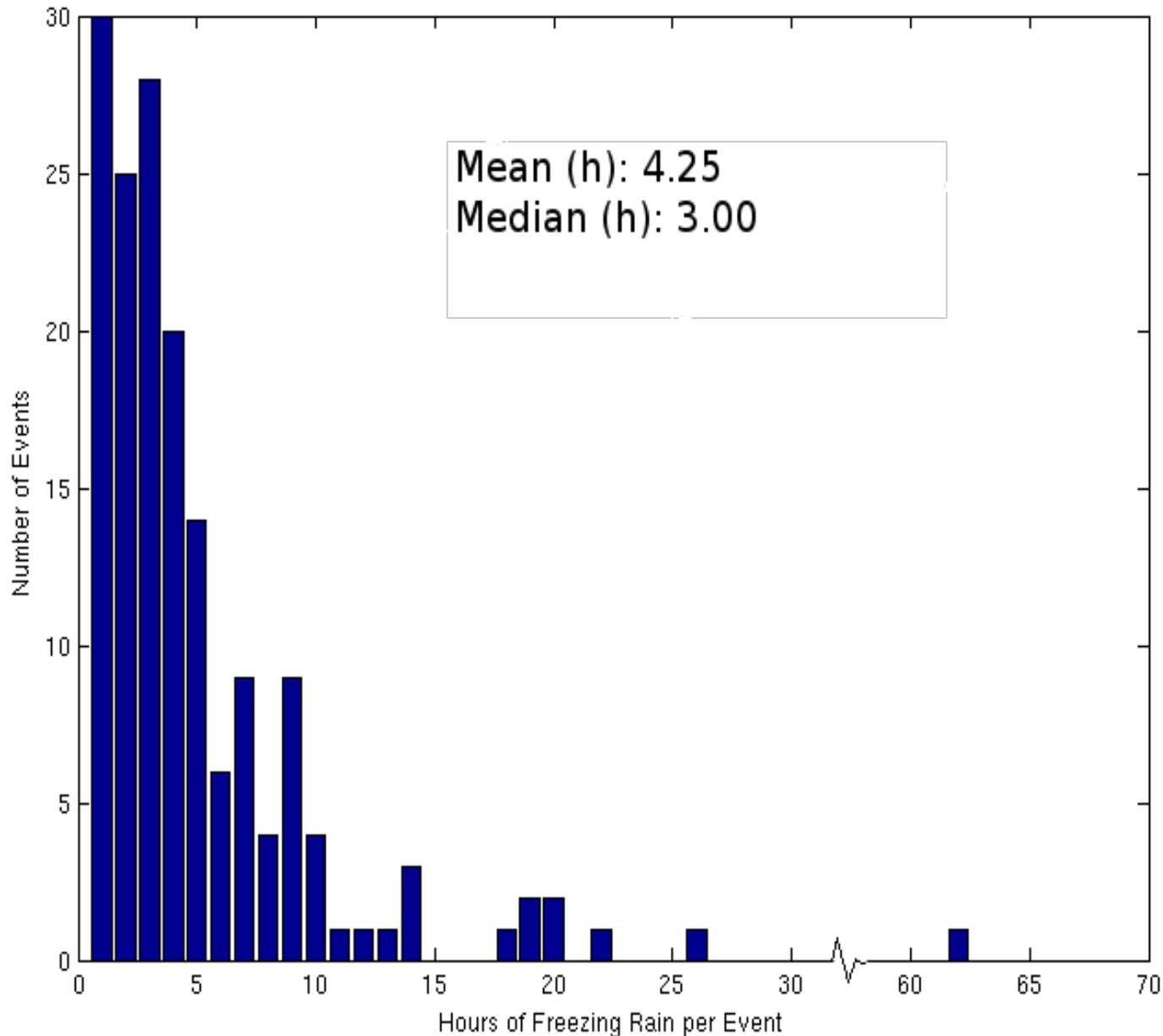
Data and Methodology

- Environment Canada hourly surface observations at Montreal, Quebec (CYUL) for the period 1979-2008
- National Centers for Environmental Prediction (NCEP) North American Regional Reanalysis (NARR) dataset



*Montreal, QC
December 2009*

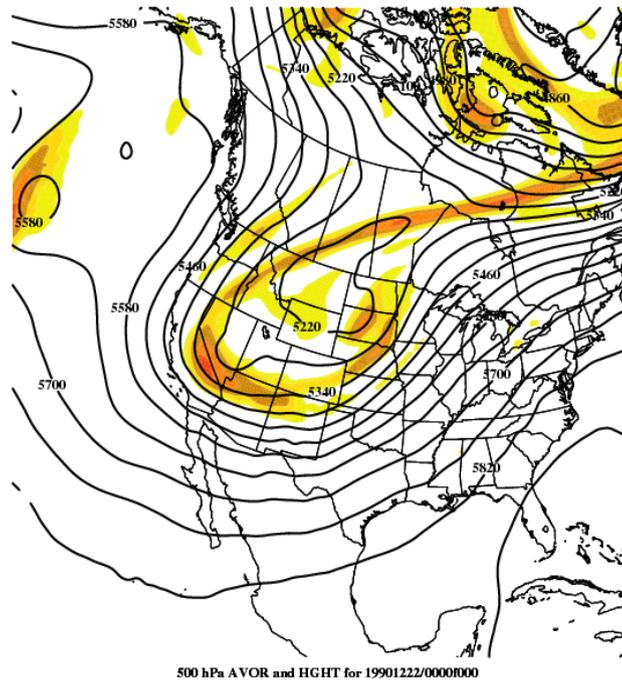
Data and Methodology



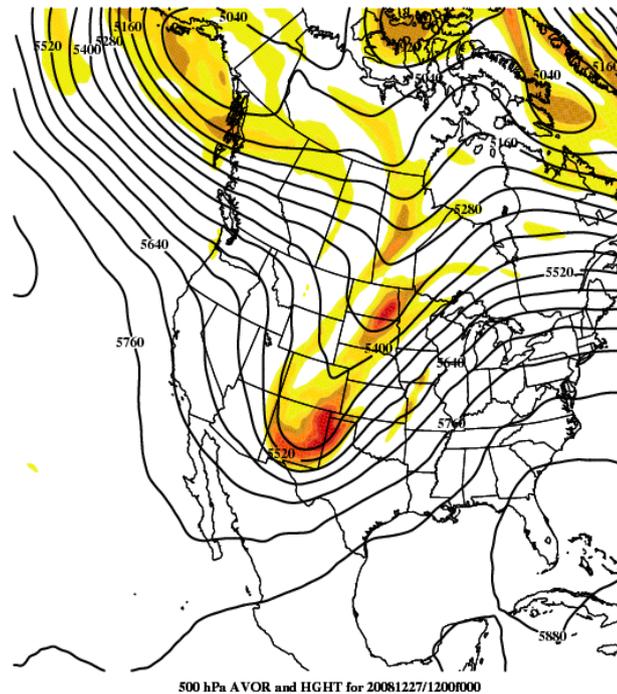
- 163 synoptically independent events
- Most are short-lived
- 98% “light” intensity
- Severe events \equiv 6 or more hours of freezing rain per event
- 46 severe events

Manual Synoptic Partitioning

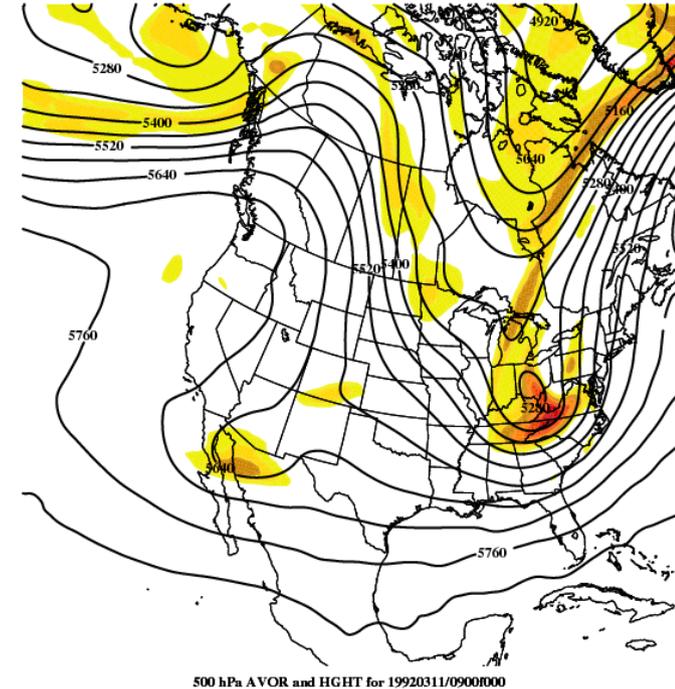
- 46 severe events were partitioned according to the location and tilt of the long-wave 500-hPa trough axis



West



Central



East

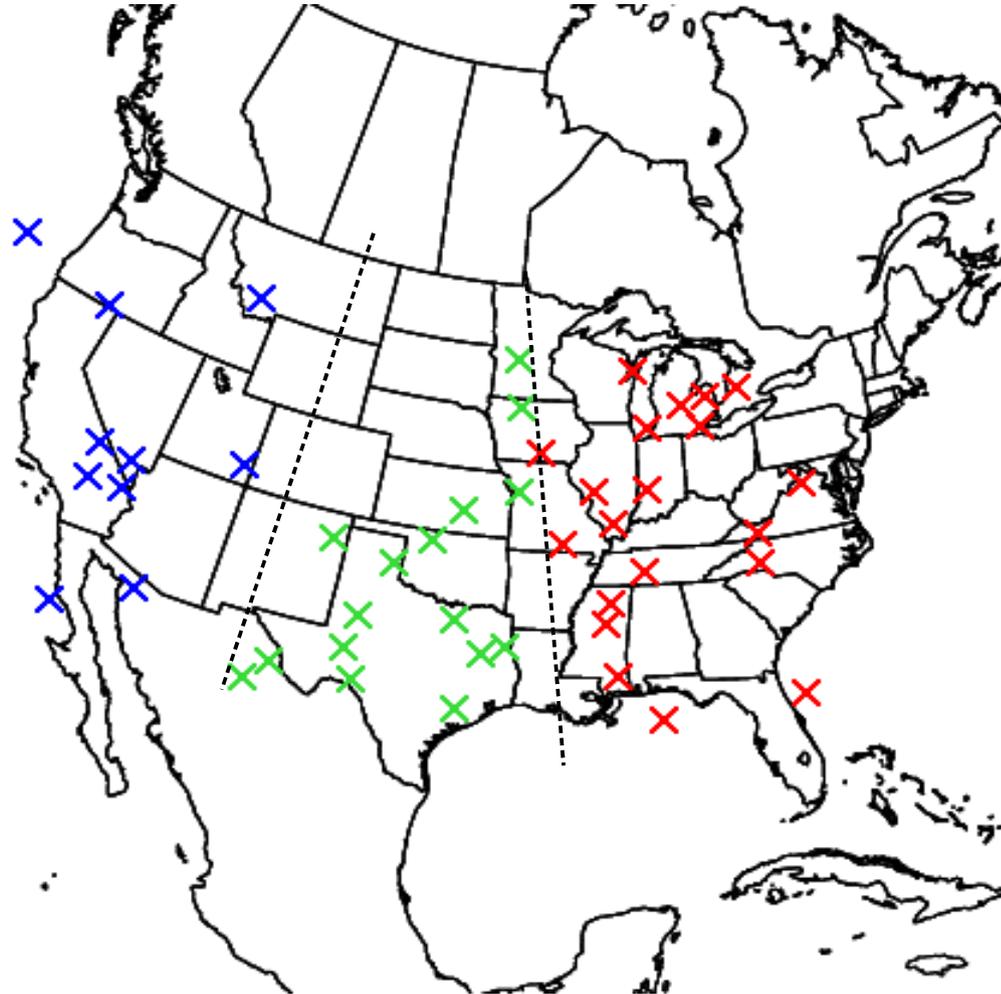
500-hPa heights (m; contoured) and absolute vorticity ($10^{-5} s^{-1}$; shaded)

Manual Synoptic Partitioning

West, n=10

Central, n=16

East, n=20



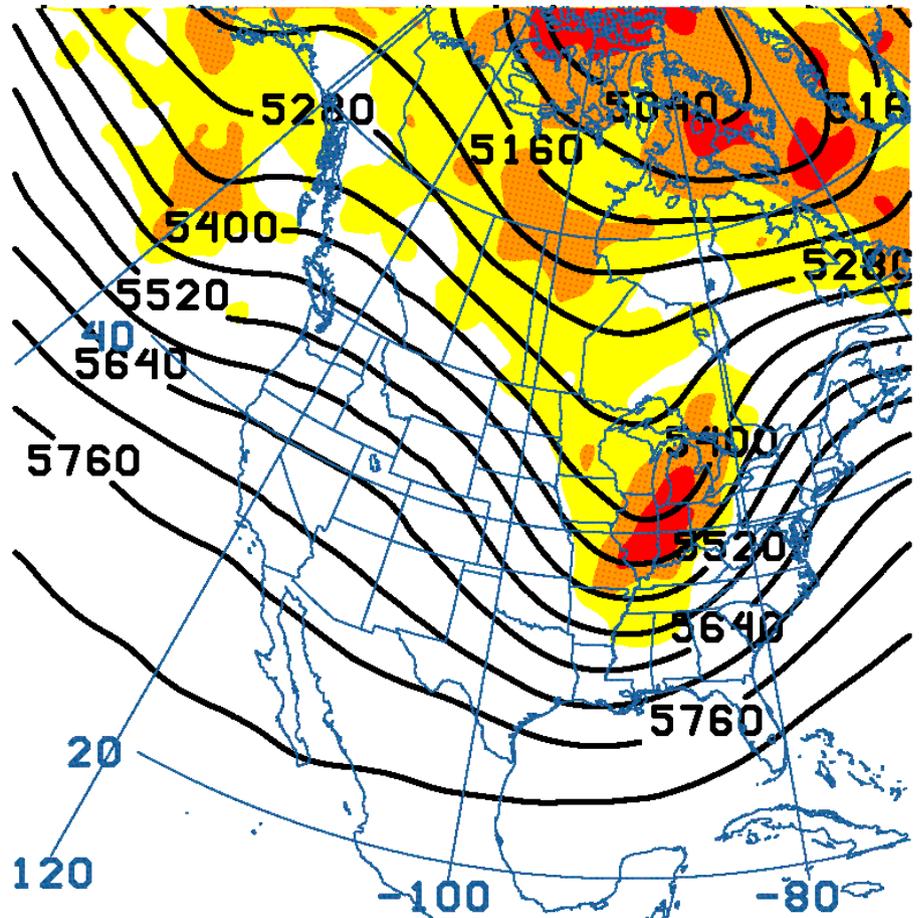
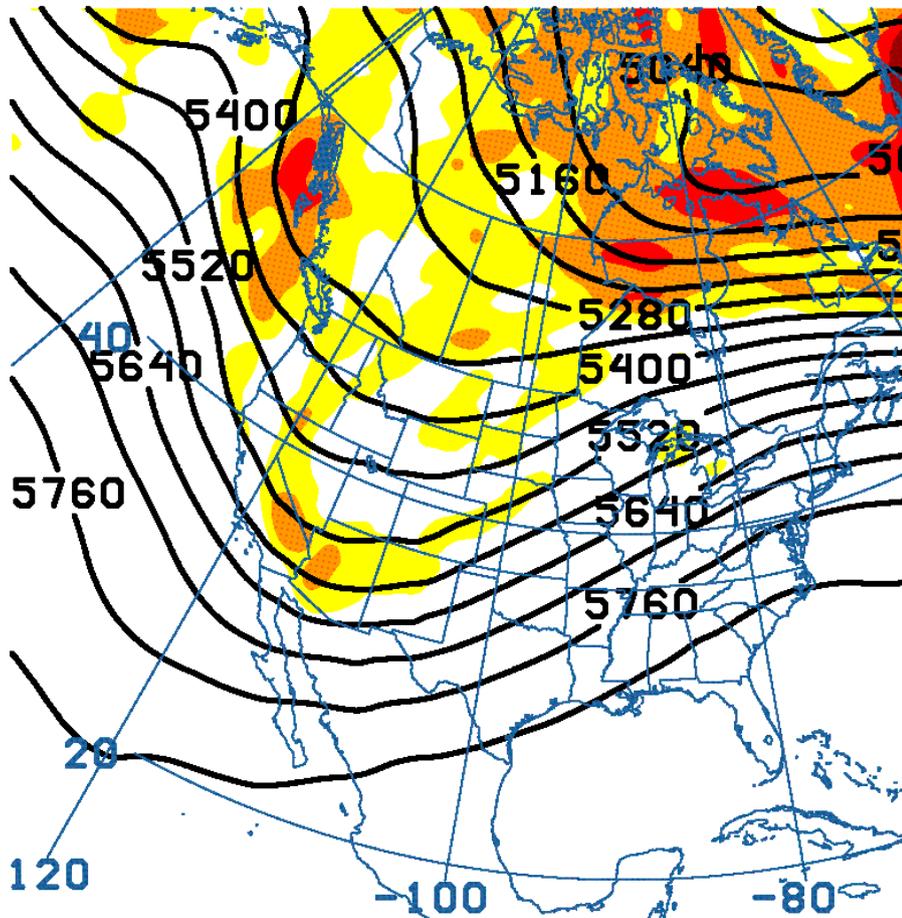
Long-wave absolute vorticity maxima for west (blue), central (green), and east (red) synoptic types

500-hPa height (m; contoured), 500-hPa absolute vorticity ($10^{-5} s^{-1}$; shaded)

t = 0 h

West, n=10

East, n=20

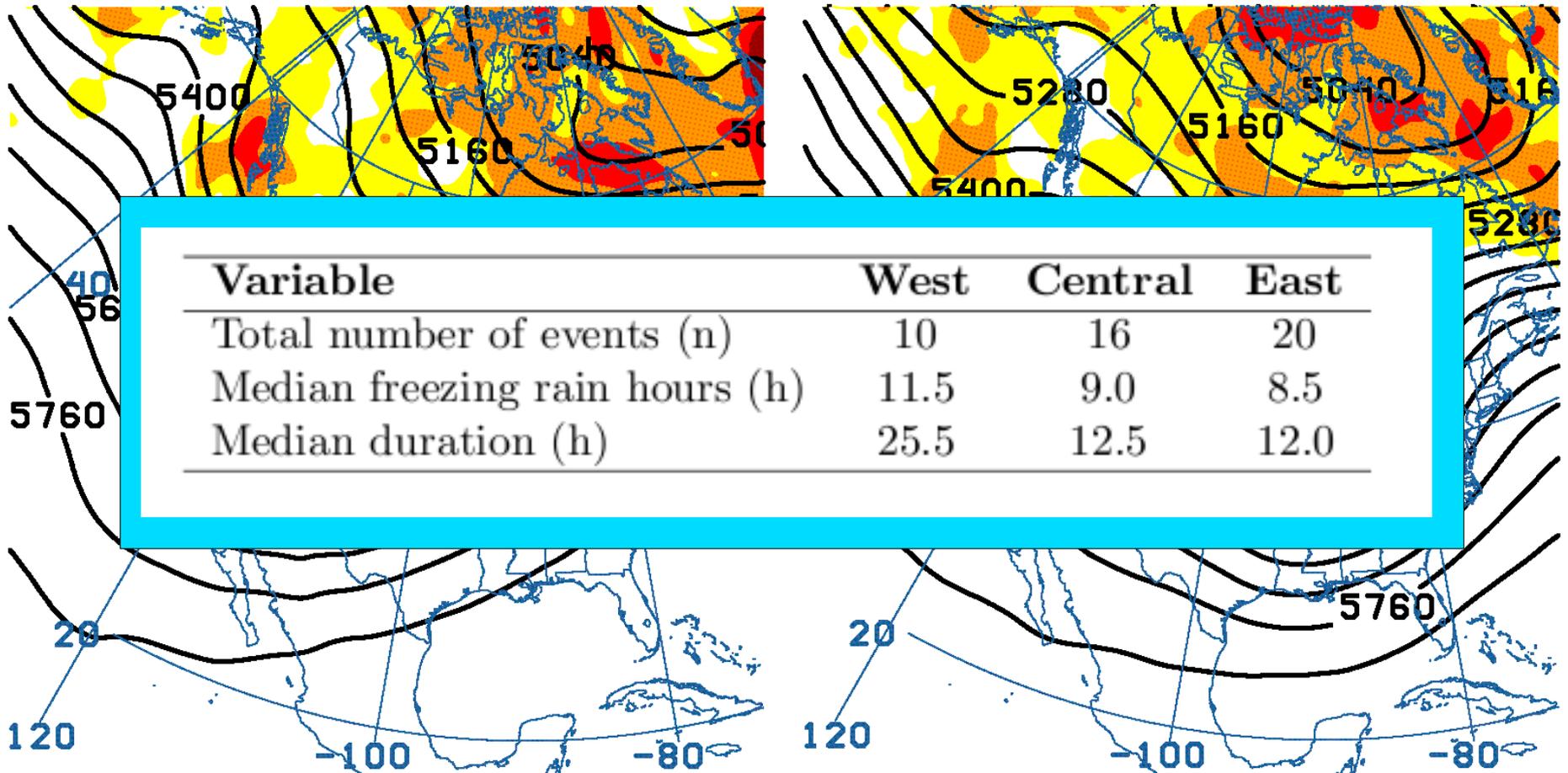


500-hPa height (m; contoured), 500-hPa absolute vorticity ($10^{-5} s^{-1}$; shaded)

t = 0 h

West, n=10

East, n=20

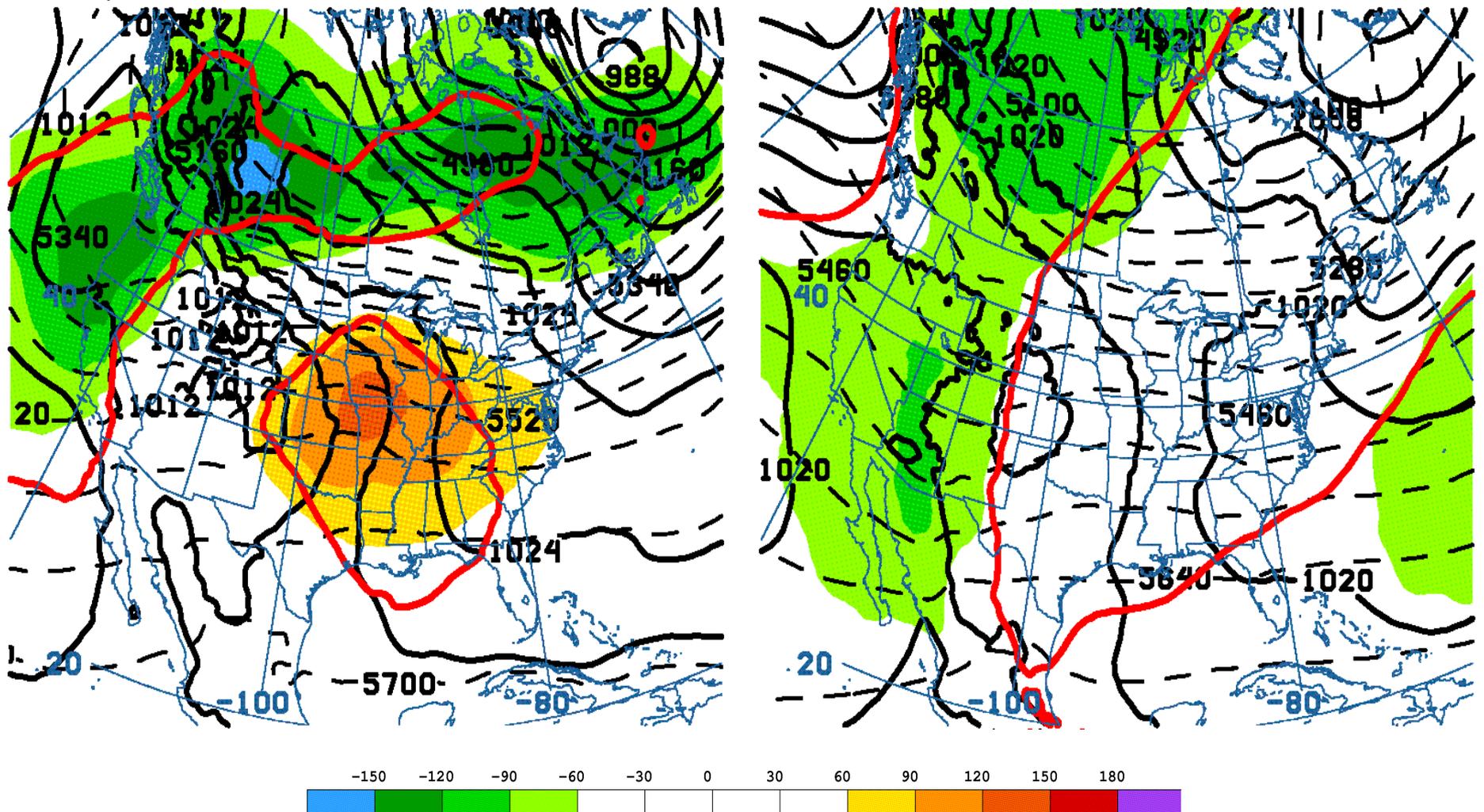


MSLP (hPa; solid), 1000-500-hPa thickness (m; dashed), thickness anomalies (m; shaded), 99% confidence interval for anomalies (red dashed)

t = -48 h

West, n=10

East, n=20

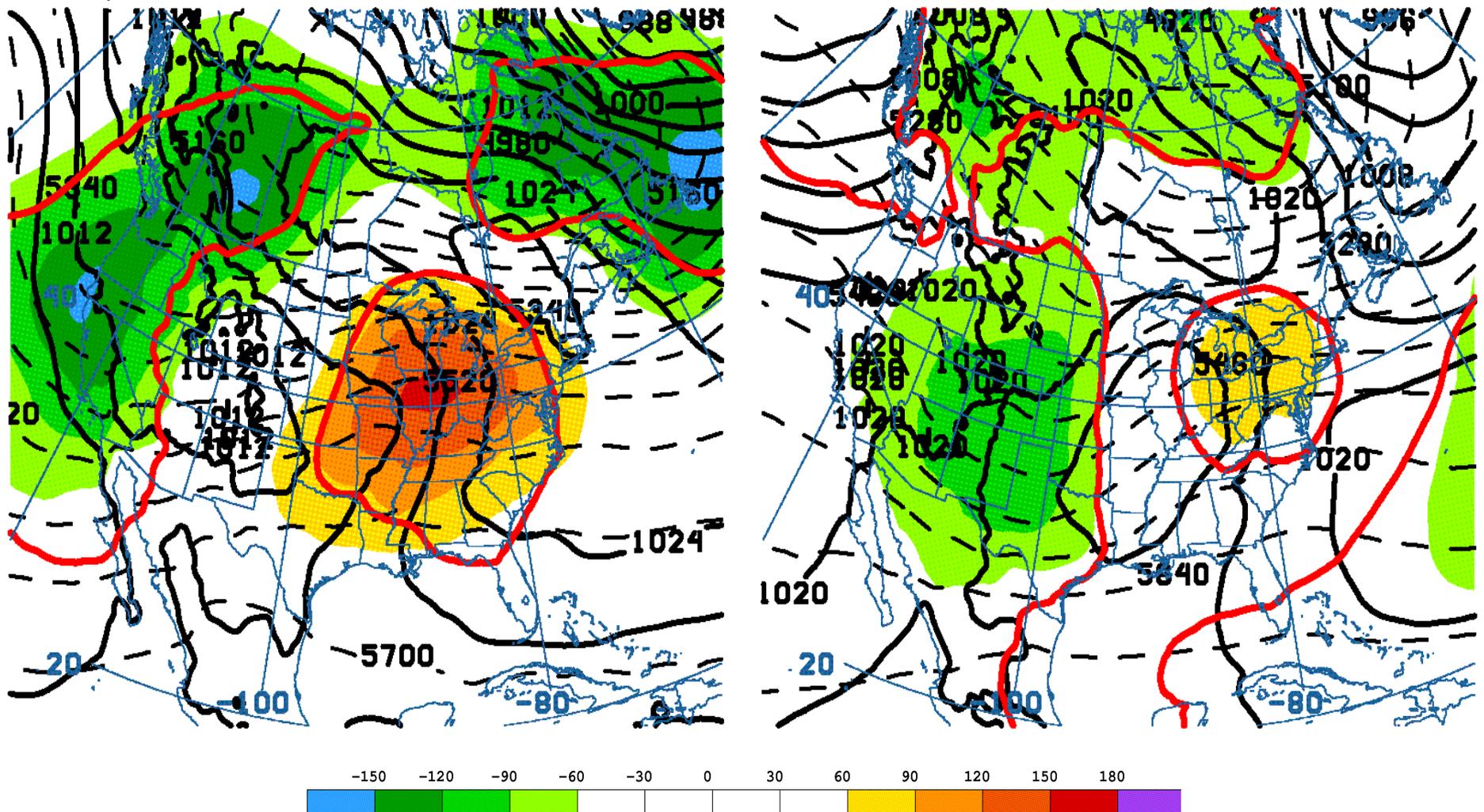


MSLP (hPa; solid), 1000-500-hPa thickness (m; dashed), thickness anomalies (m; shaded), 99% confidence interval for anomalies (red dashed)

t = -24 h

West, n=10

East, n=20

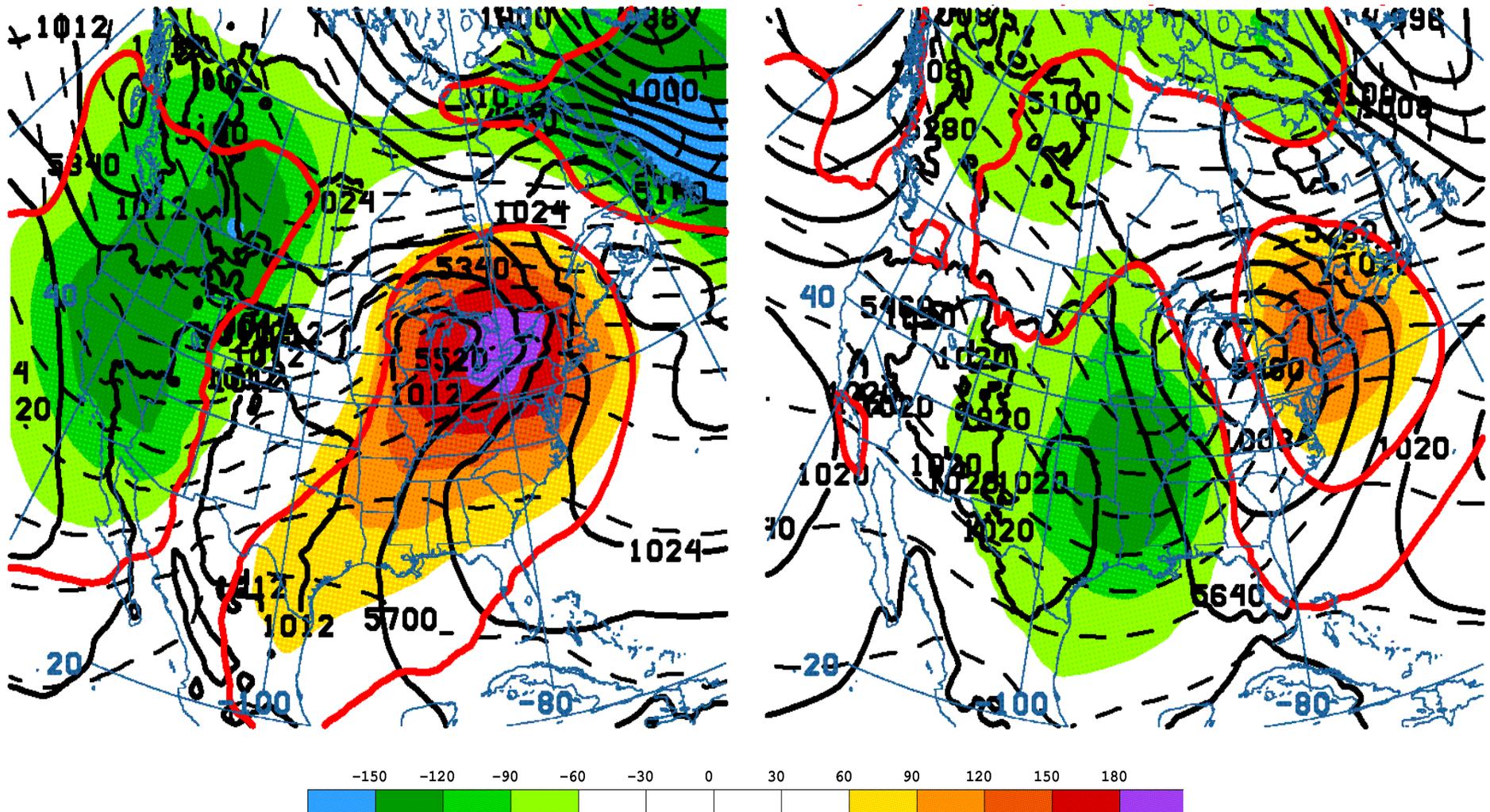


MSLP (hPa; solid), 1000-500-hPa thickness (m; dashed), thickness anomalies (m; shaded), 99% confidence interval for anomalies (red dashed)

t = 0 h

West, n=10

East, n=20

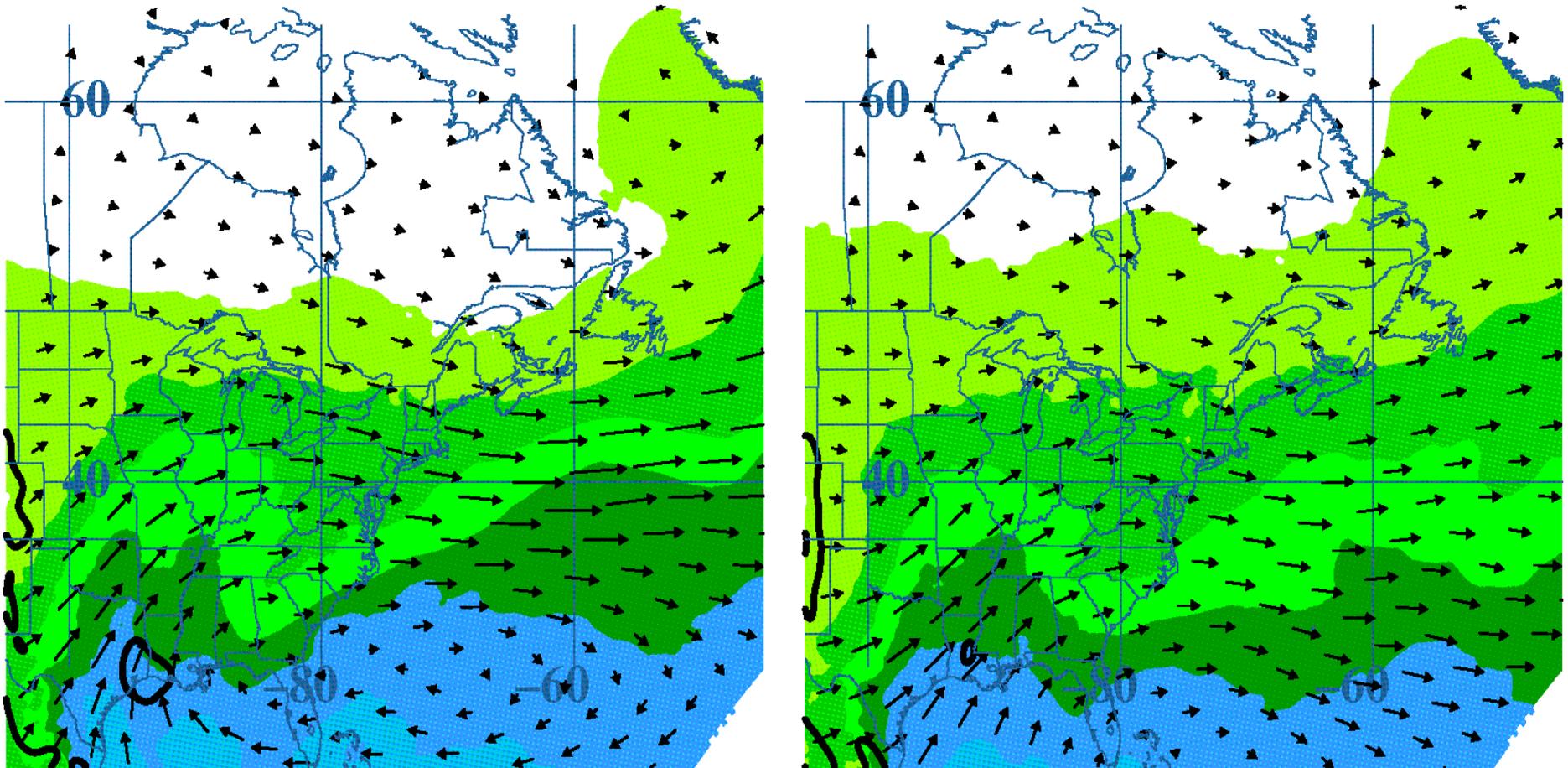


Precipitable water (mm; shaded), water vapour transport ($\text{kg m}^{-1} \text{s}^{-1}$; arrows), water vapour transport convergence ($-2 \times 10^{-7} \text{ kg m}^{-1} \text{s}^{-1}$; contoured)

t = -48 h

West, n=10

East, n=20

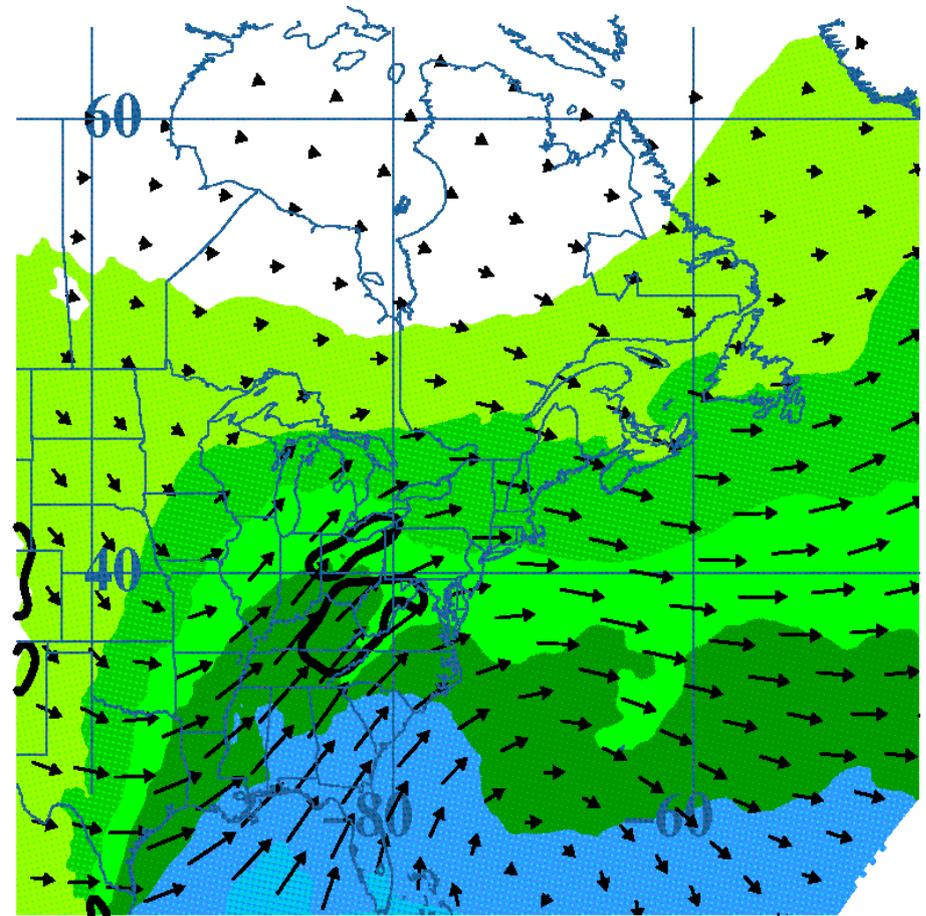
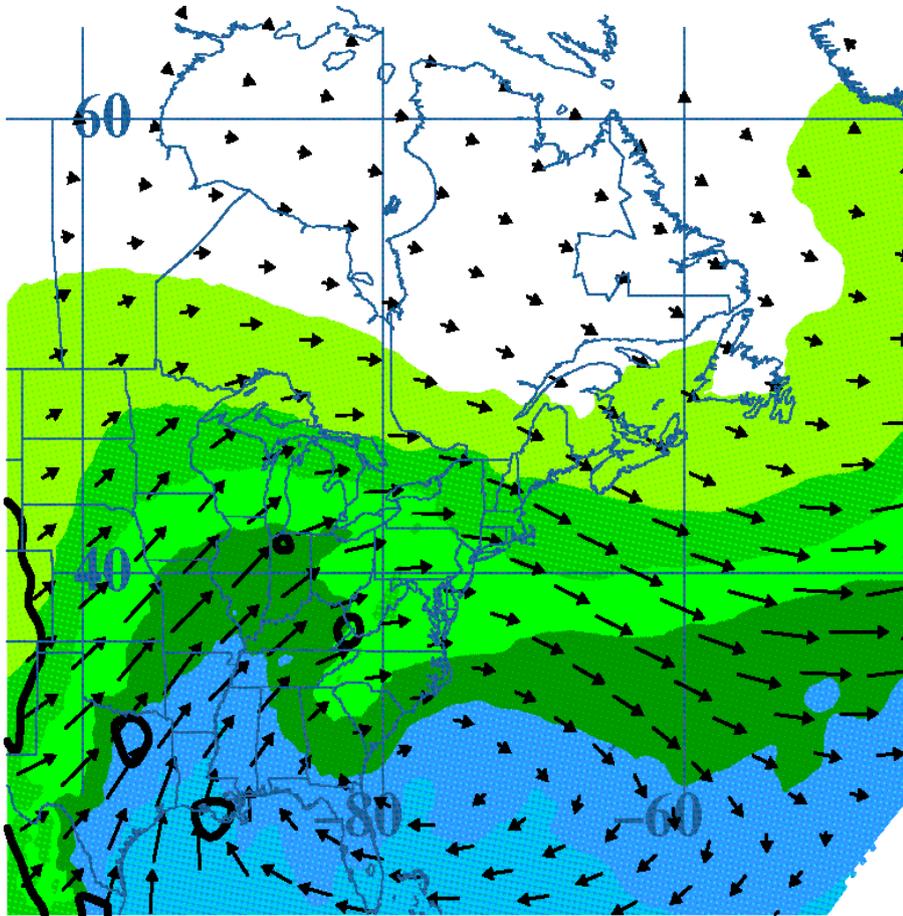


Precipitable water (mm; shaded), water vapour transport ($\text{kg m}^{-1} \text{s}^{-1}$; arrows), water vapour transport convergence ($-2 \times 10^{-7} \text{ kg m}^{-1} \text{s}^{-1}$; contoured)

t = -24 h

West, n=10

East, n=20

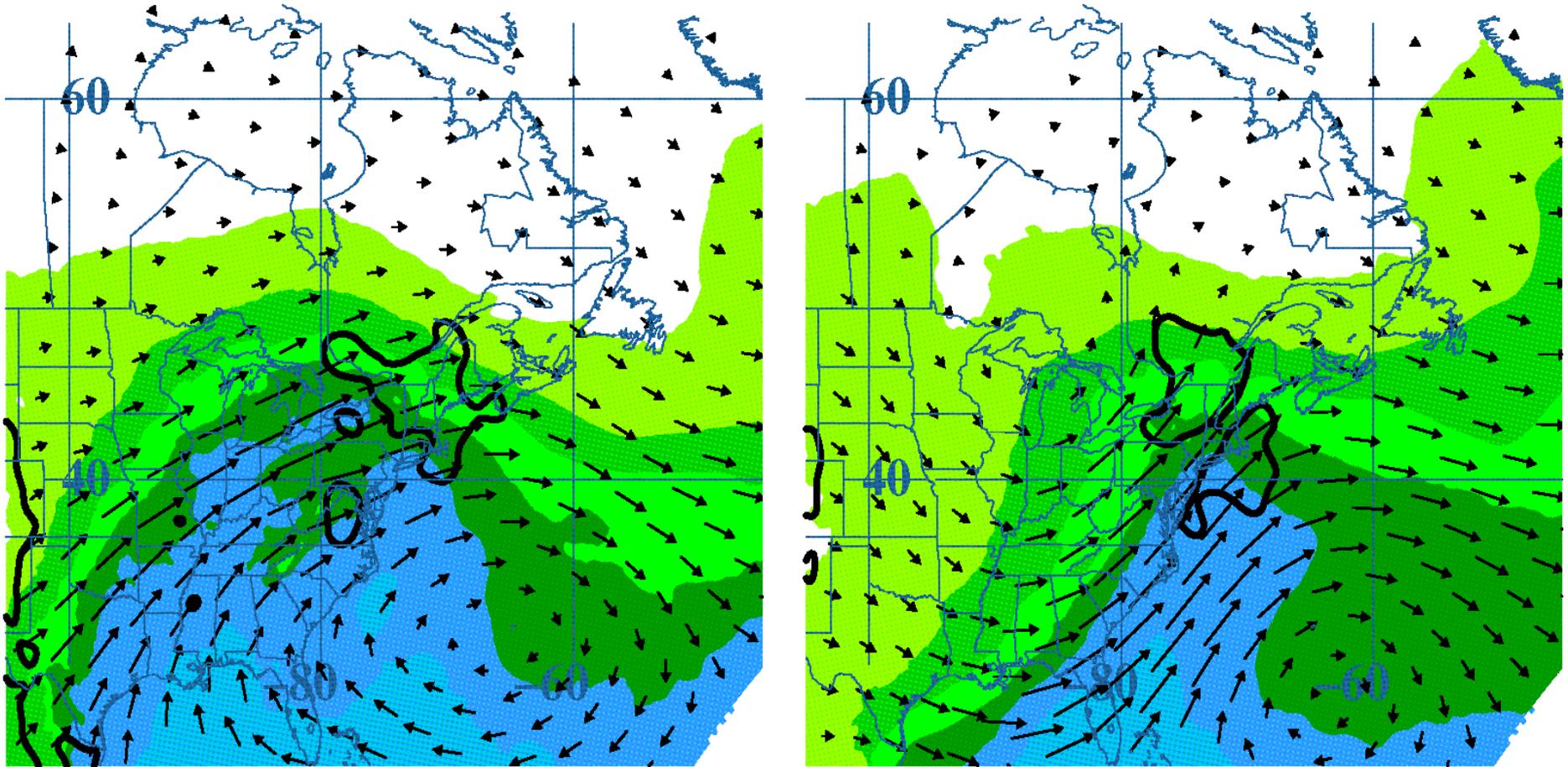


Precipitable water (mm; shaded), water vapour transport ($\text{kg m}^{-1} \text{s}^{-1}$; arrows), water vapour transport convergence ($-2 \times 10^{-7} \text{ kg m}^{-1} \text{s}^{-1}$; contoured)

t = 0 h

West, n=10

East, n=20

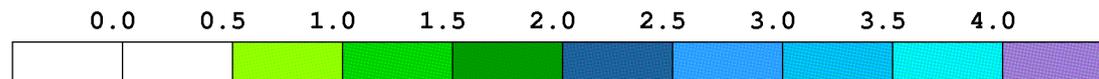
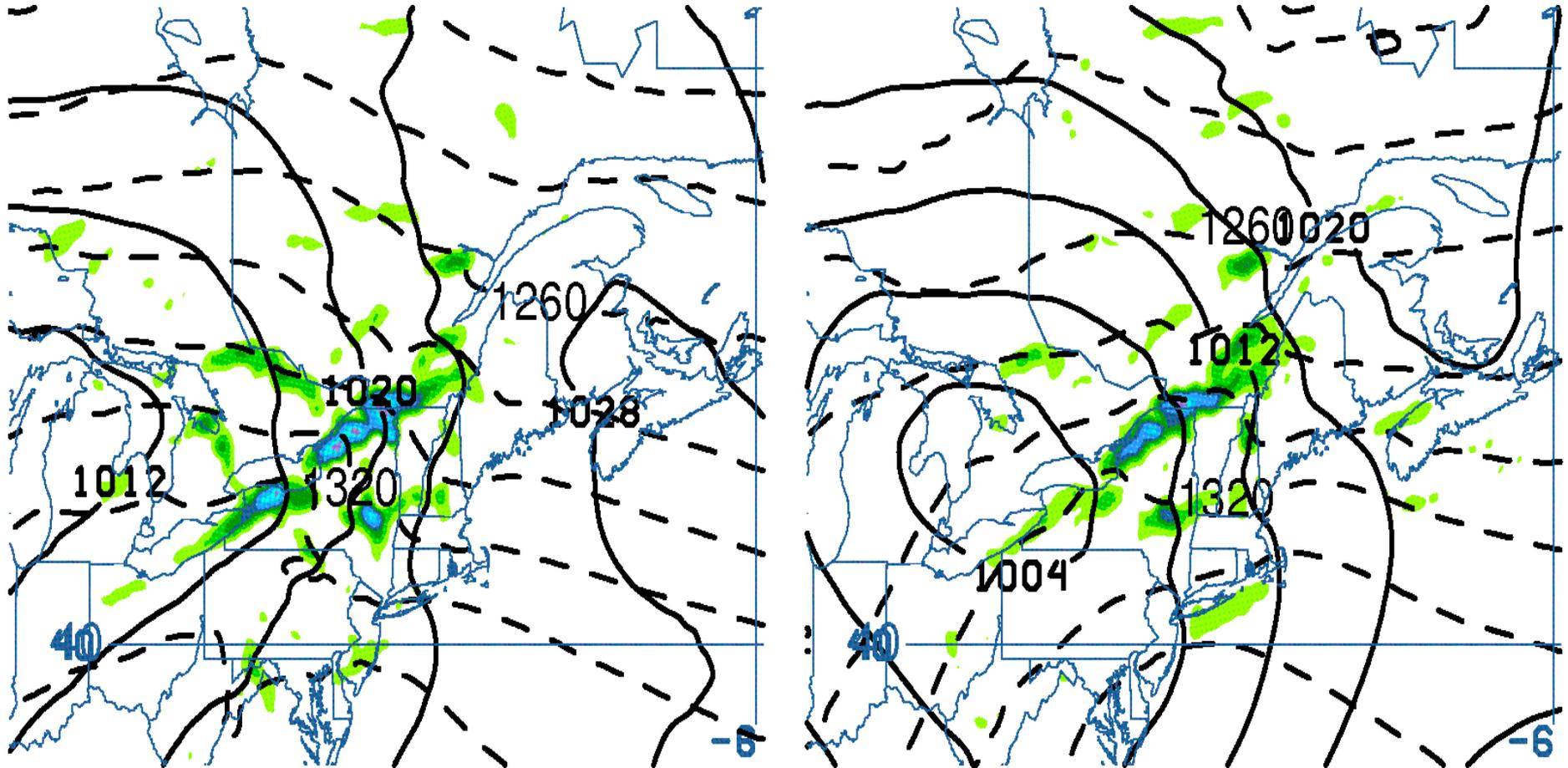


*MSLP (hPa; solid), 1000-850-hPa thickness (m; dashed),
1000-850-hPa frontogenesis ($K (100km)^{-1} (3h)^{-1}$; shaded)*

t = 0 h

West, n=10

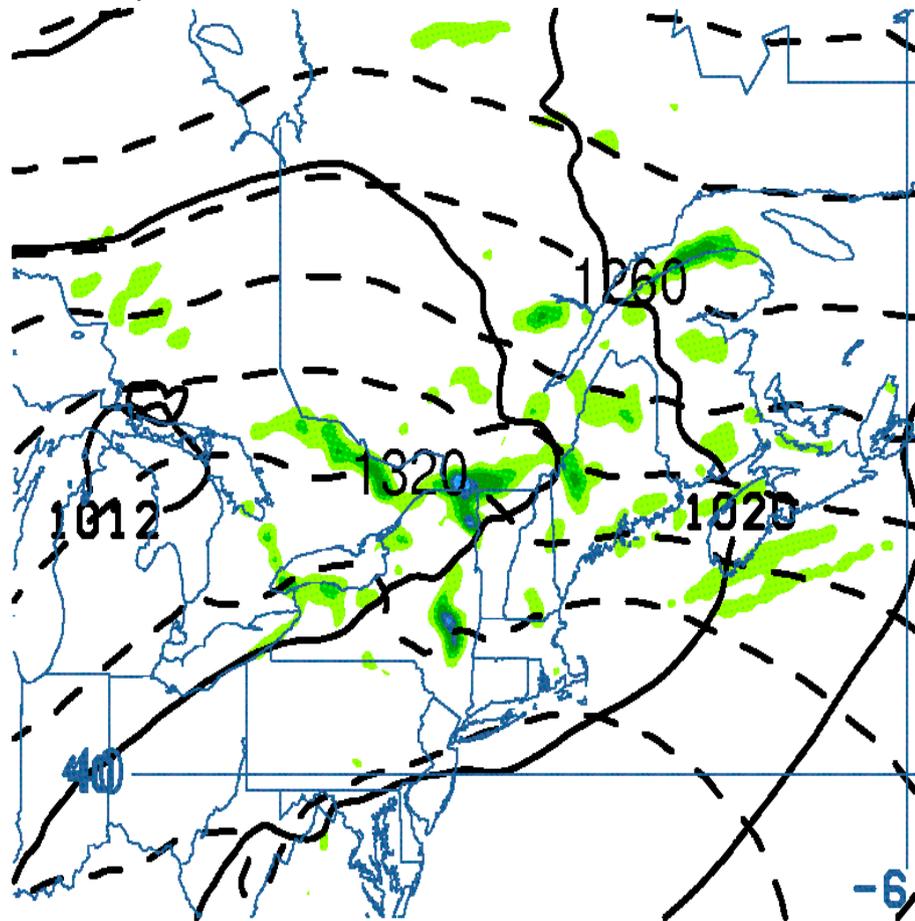
East, n=20



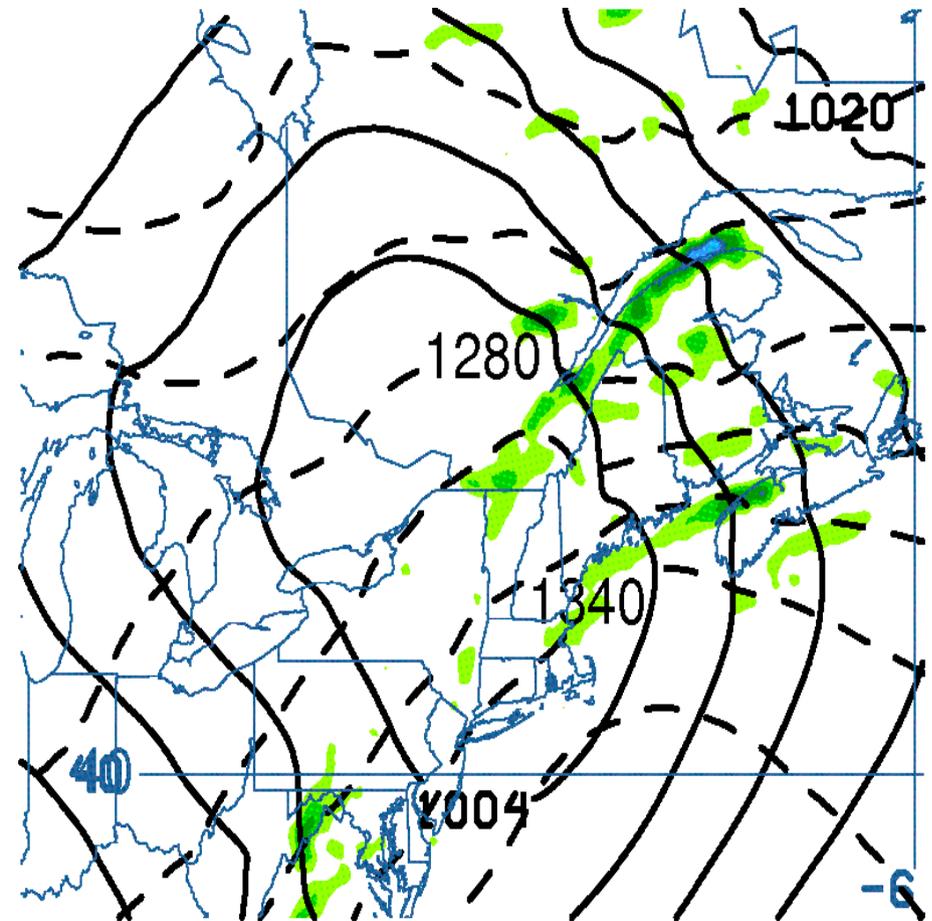
*MSLP (hPa; solid), 1000-850-hPa thickness (m; dashed),
1000-850-hPa frontogenesis ($K (100km)^{-1} (3h)^{-1}$; shaded)*

t = +12 h

West, n=10



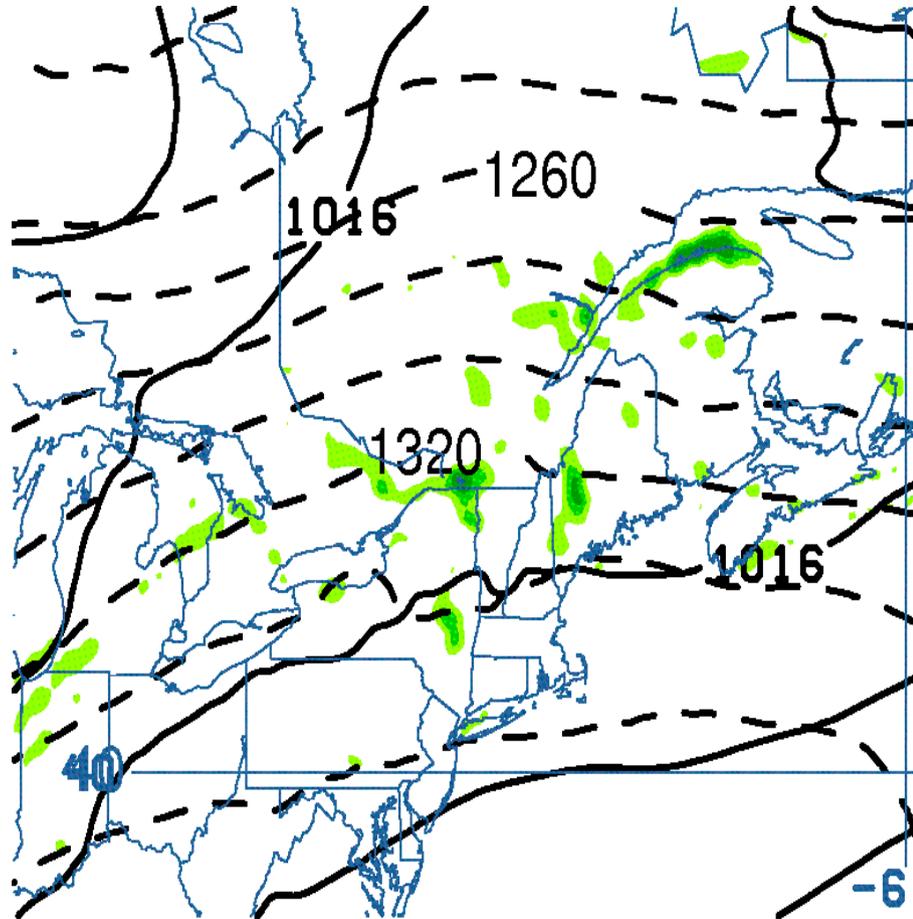
East, n=20



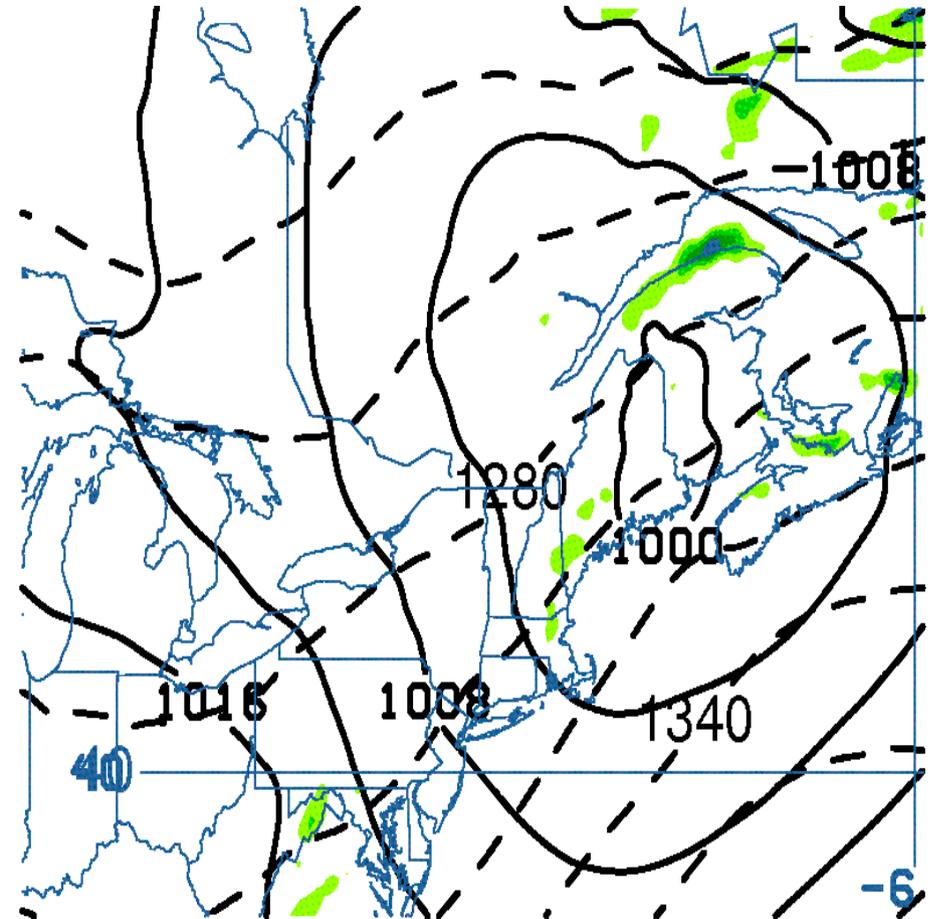
*MSLP (hPa; solid), 1000-850-hPa thickness (m; dashed),
1000-850-hPa frontogenesis ($K (100km)^{-1} (3h)^{-1}$; shaded)*

t = +24 h

West, n=10



East, n=20



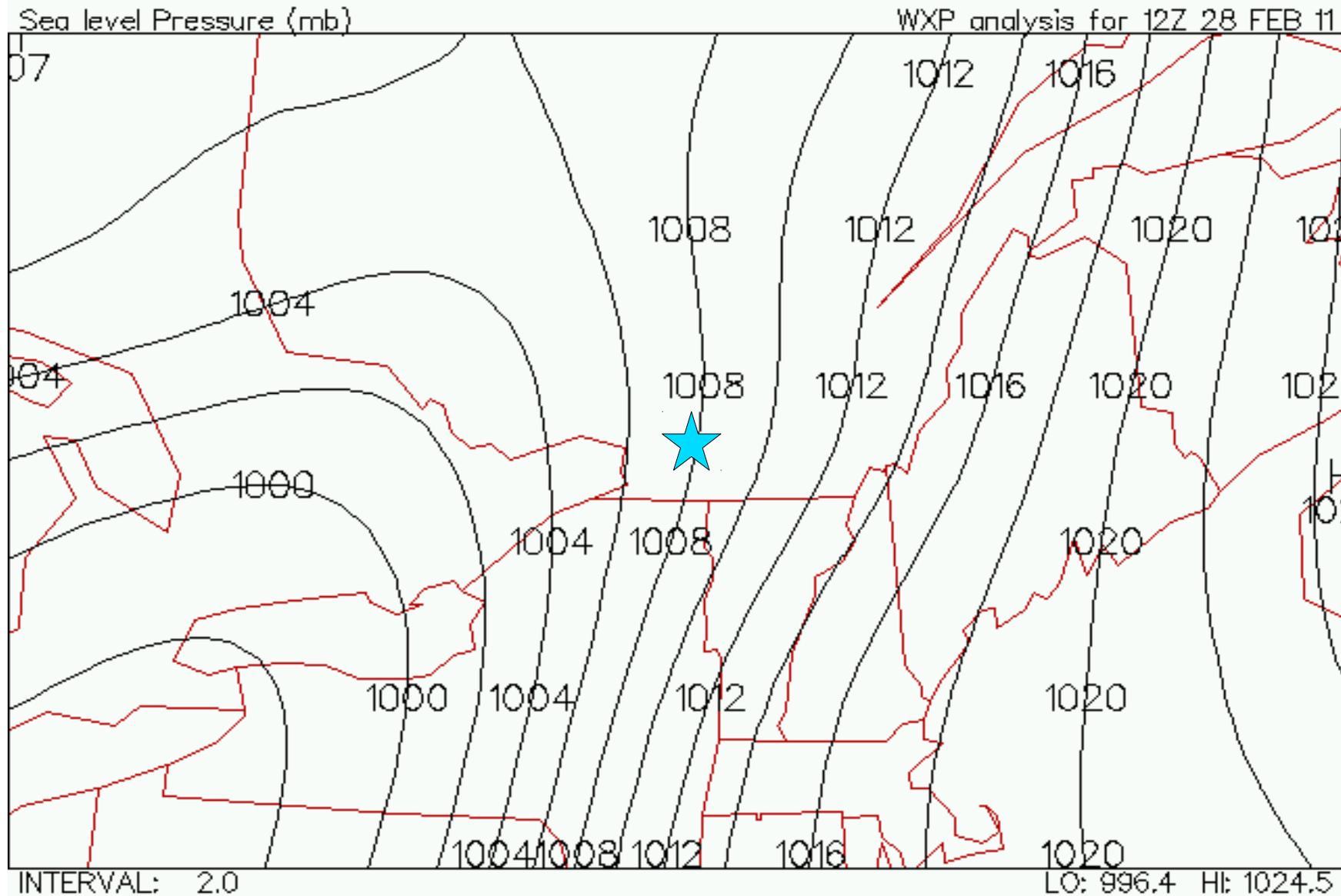
Composite Analysis Conclusions

- WEST composite
 - Long upper-level wavelength: slower eastward movement
 - Southwesterly geostrophic flow: moisture from the Gulf of Mexico
 - Valley-enhanced frontogenesis
- EAST composite
 - Short upper-level wavelength: faster eastward movement
 - Southeasterly geostrophic flow: moisture from the East Coast
 - Passage of a midlatitude cyclone

*For sustained freezing rain, there must be a **pressure gradient force** aligned with the St. Lawrence River Valley – whether it results from a cyclone, anticyclone, or couplet.*

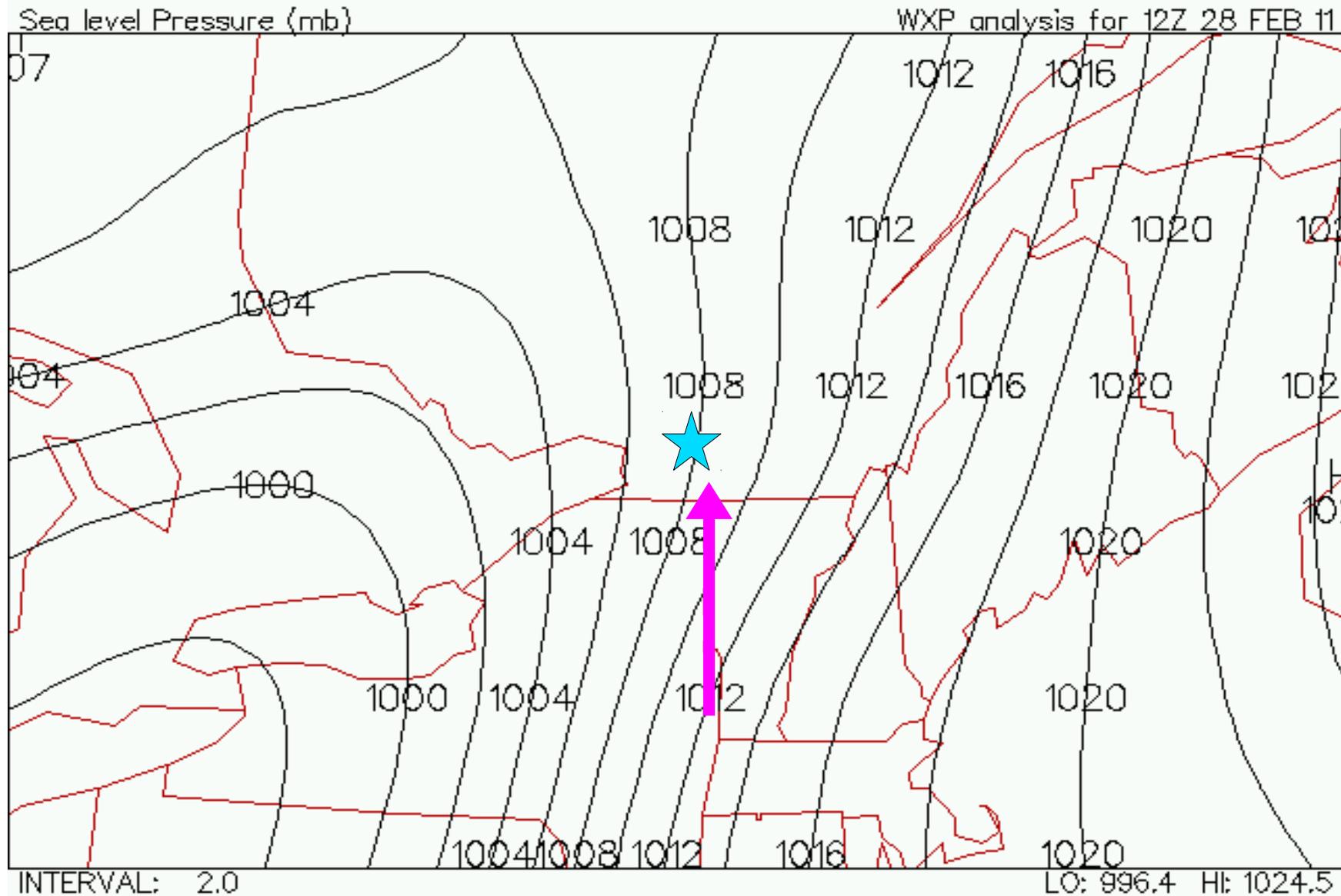
Just one caveat...

▼ Plymouth State Weather Center ▼



Southerly channeling!

▼ Plymouth State Weather Center ▼

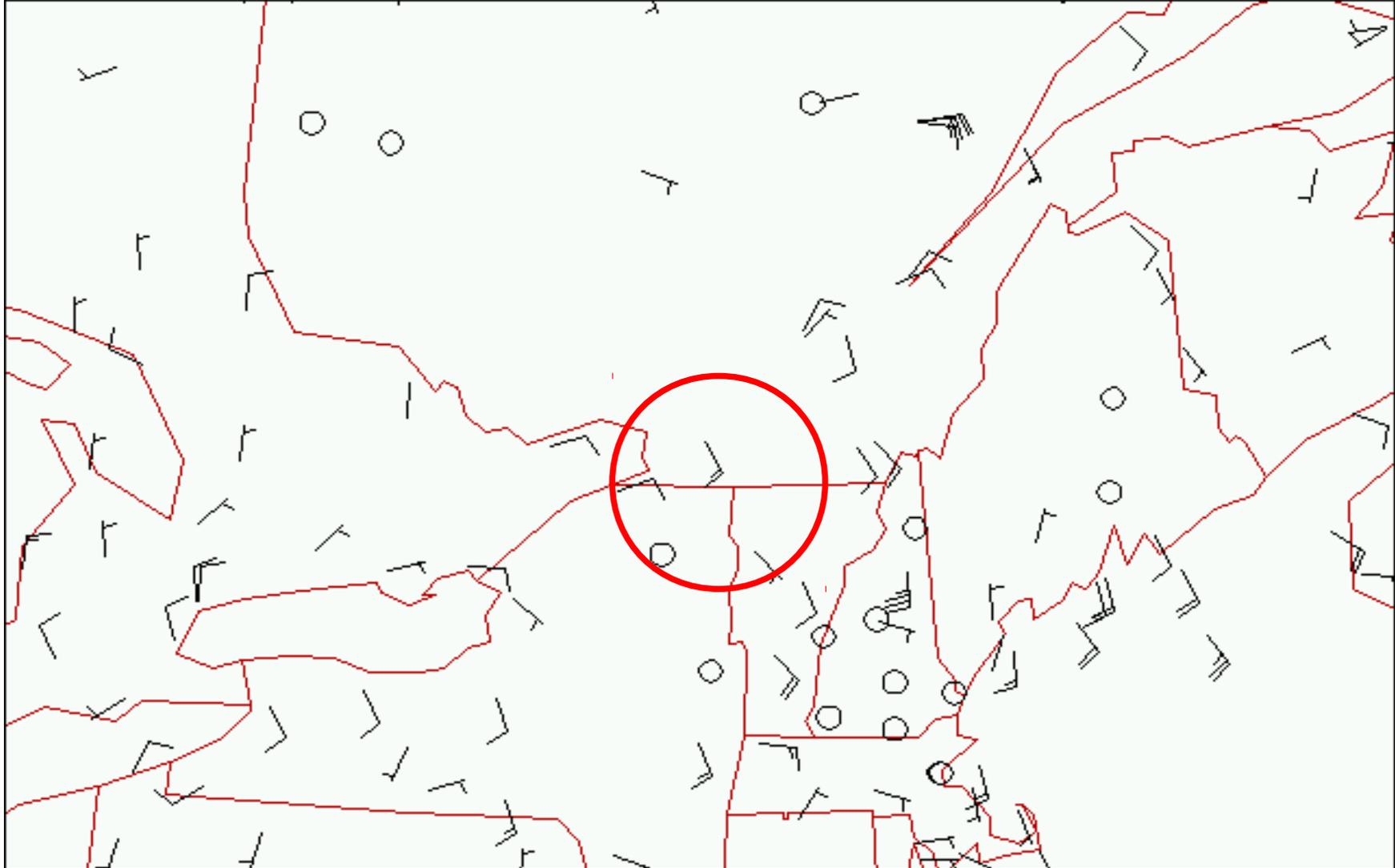


Southerly channeling!

▼ Plymouth State Weather Center ▼

Surface Winds (knt)

Analysis for 12Z 28 FEB 11



Southerly channeling!

▼ Plymouth State Weather Center ▼

Surface Winds (knt)

Analysis for 12Z 28 FEB 11

```
CYUL 280900Z 03009KT 15SM OVC046 M05/M08 A2982 RMK SC8 SLP099
CYUL 281000Z 02008KT 15SM OVC047 M04/M07 A2980 RMK SC8 SLP093
CYUL 281100Z 15009KT 15SM -SN OVC038 01/M03 A2977 RMK SC8 SLP082
CYUL 281200Z 15016KT 2SM -RASN SCT011 OVC022 00/M02 A2970 RMK
SF3SC5 PRESFR SLP059
CYUL 281300Z 15013G21KT 3SM -RASN SCT010 OVC022 00/M02 A2965 RMK
SF3SC5 SLP043
```

